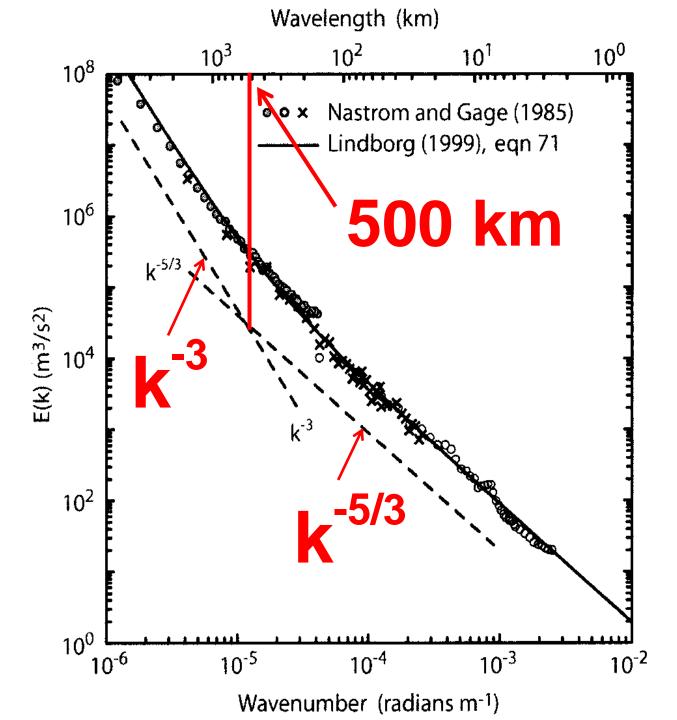
Betydelsen av sammanväxten av data-assimilation och ensemble-teknik för Nowcasting

Åke Johansson SMHI

Prediktabilitet



Allerstädes närvarande

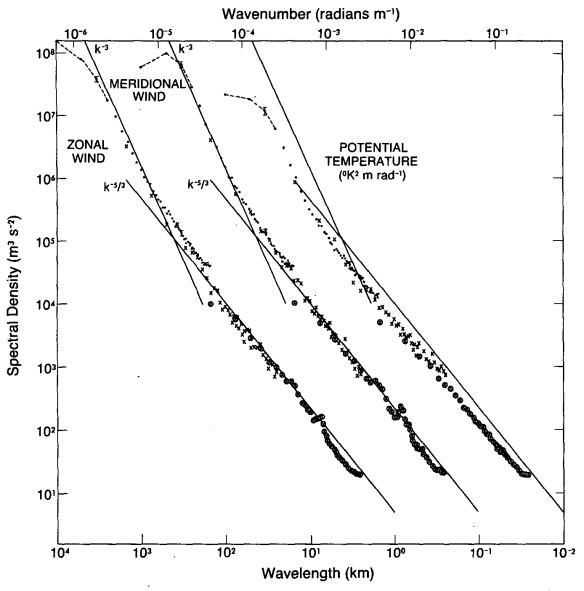
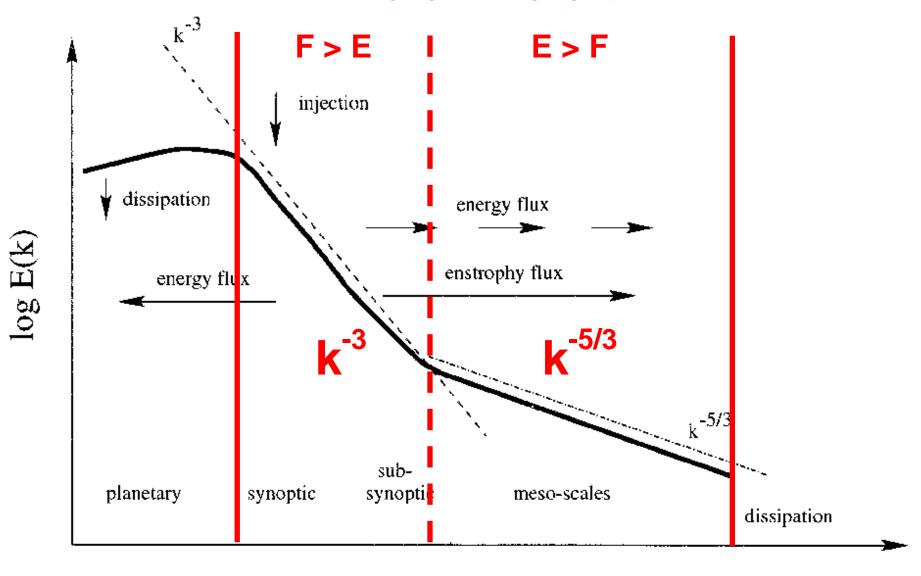


Fig. 3. Variance power spectra of wind and potential temperature near the tropopause from GASP aircraft data. The spectra for meridional wind and temperature are shifted one and two decades to the right, respectively; lines with slopes -3 and $-\frac{5}{3}$ are entered at the same relative coordinates for each variable for comparison.

F and E invariant



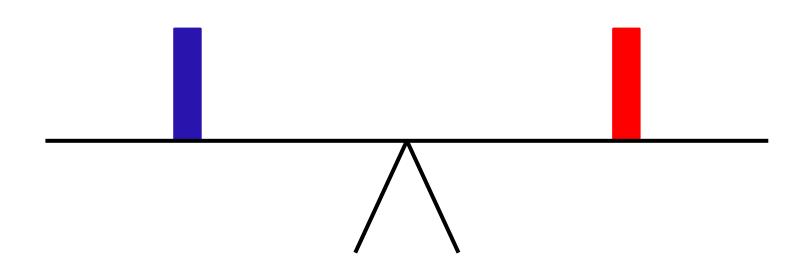
log k

Synoptic scale flow Conservation laws

$$E = \frac{1}{2} \overline{V \cdot V}$$
 Energy

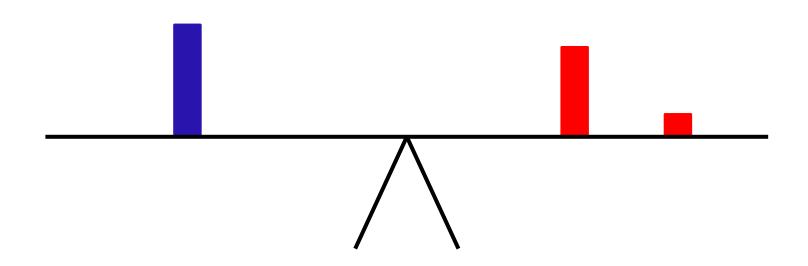
$$F = \frac{1}{2} \overline{\zeta \cdot \zeta}$$
 Enstrophy

Energy and Enstrophy Conservation



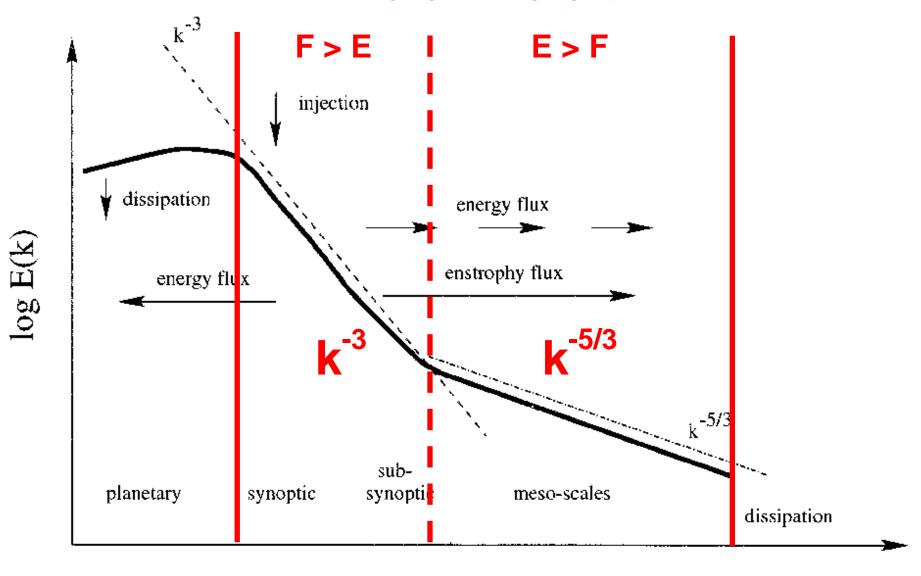
Mechanical analogy

Energy and Enstrophy Conservation



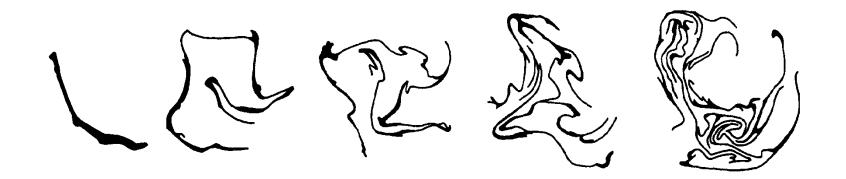
Mechanical analogy

F and E invariant



log k

Enstrophy Cascade



- Moves to smaller scale of motion
- Through differential advection
- Finally destroyed by microscale turbulence

Energy Cascade

Moves to smaller scale of motion

- Through vortex stretching
- Finally destroyed by microscale turbulence

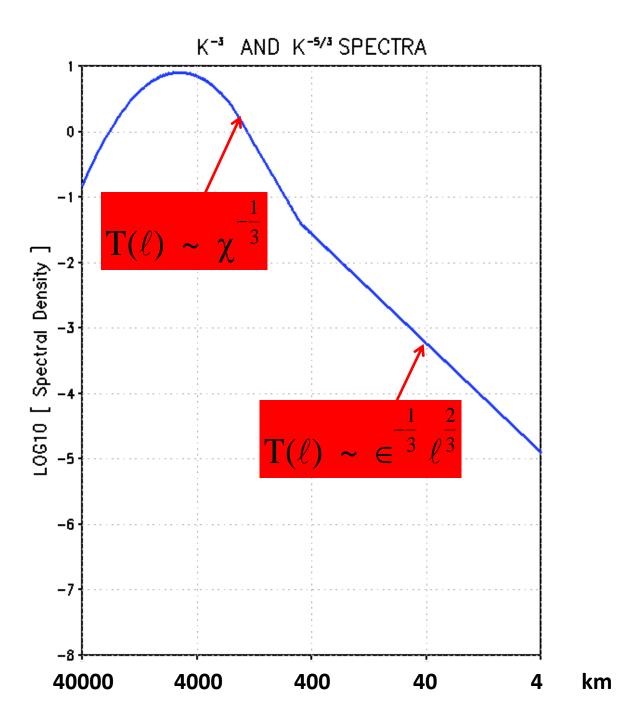
Big whirls have little whirls That feed on their velocity And little whirls have lesser whirls And so on to viscosity

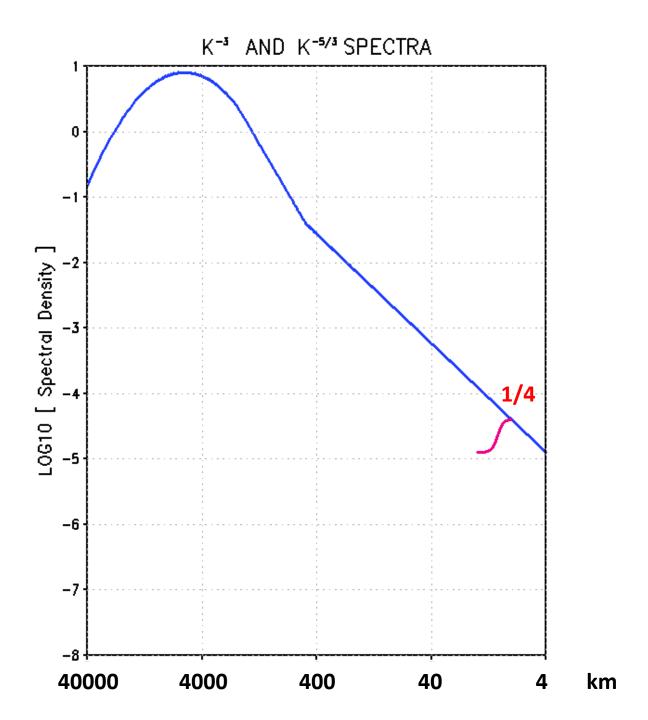
Lewis Fry Richardson

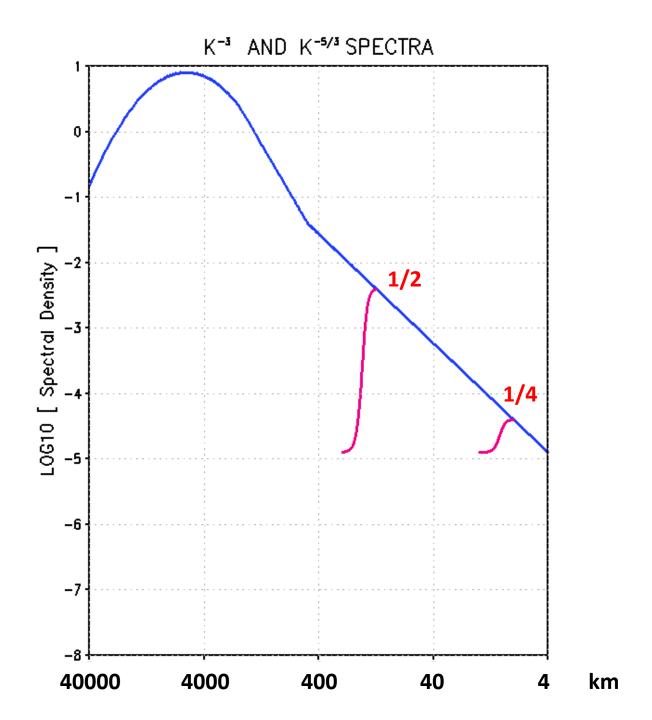
Cascade rates are intimately connected with

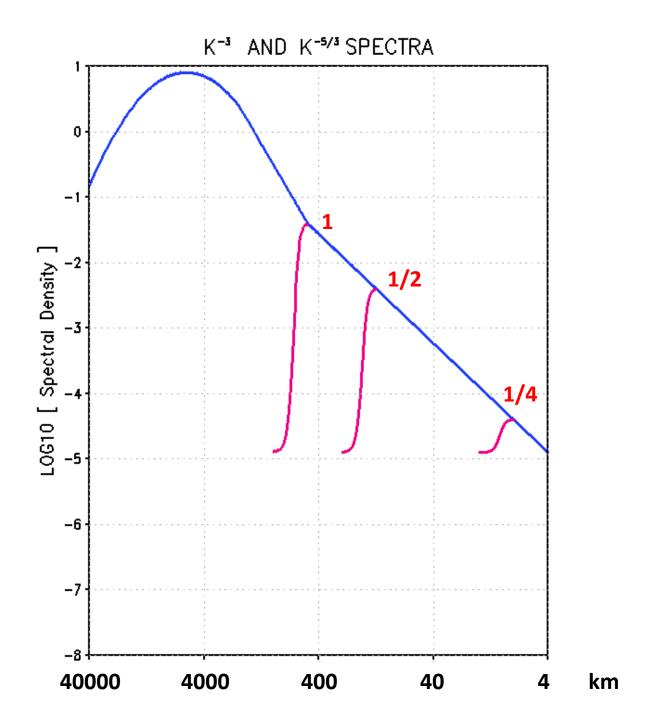
(i) Eddy turnover times in a turbulent fluid

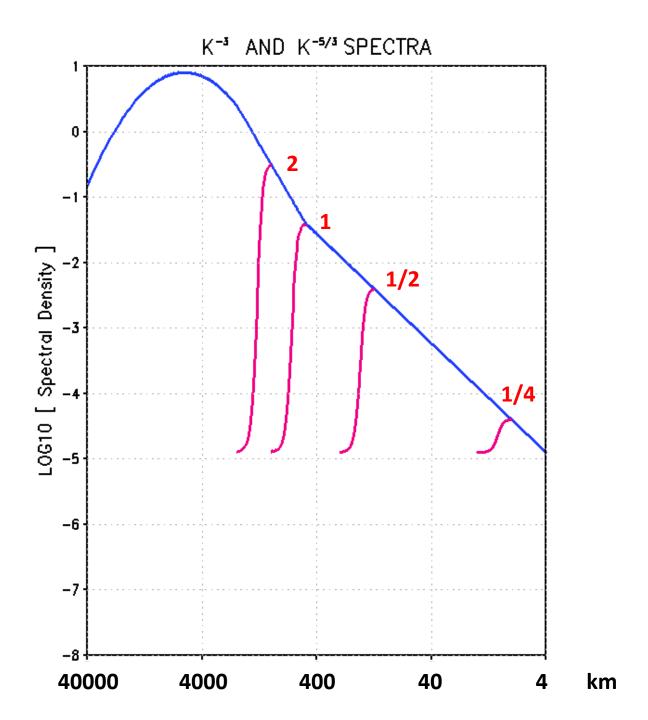
(ii) Spectral slope of the energy spectrum

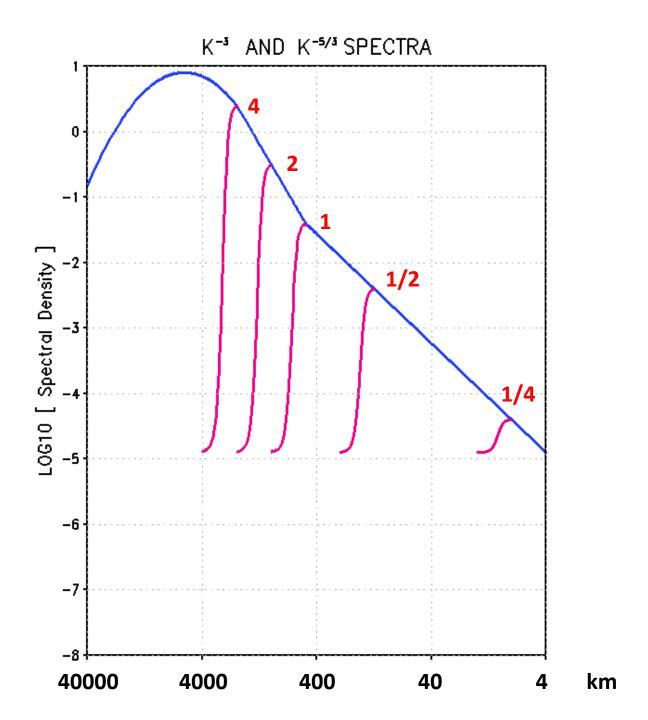


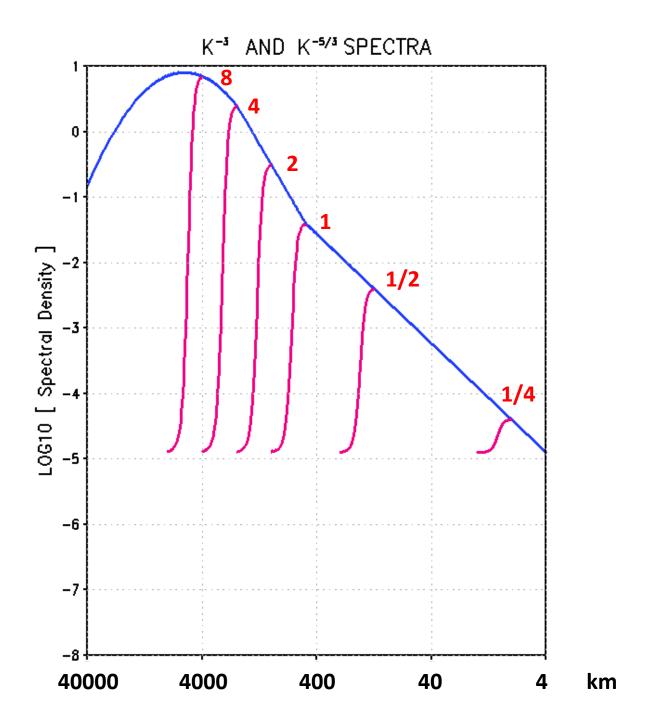




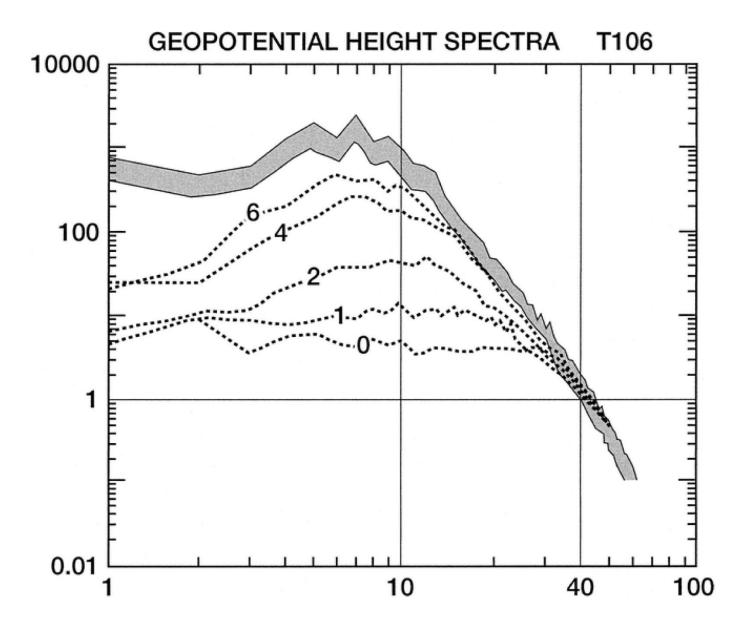








Realistic Error Spectrum



Tribbia and Baumhefner 2004

Three types of Error Growth

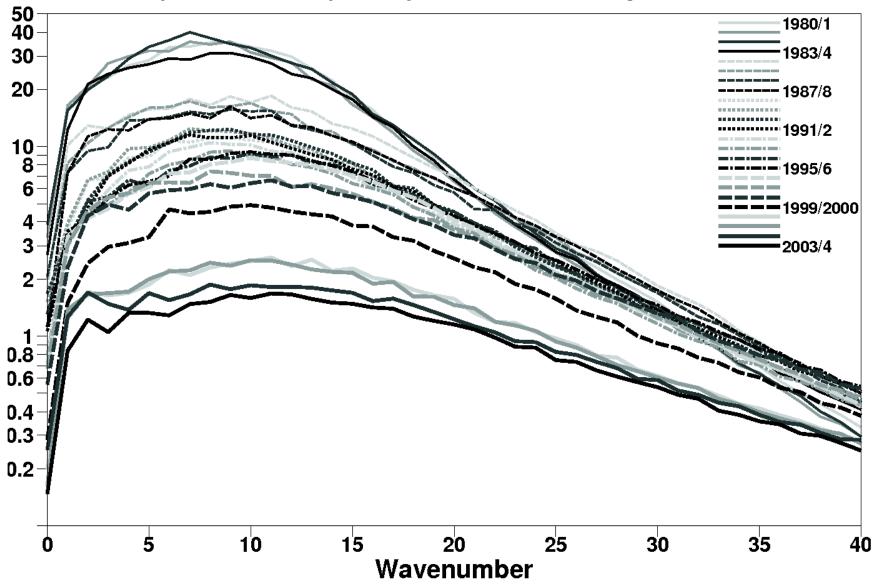
1. Inverse cascade

2. Baroclinic instability

3. Advection

Observed Improvements in Initial Error

Spectra of mean square day-1 error of 500hPa height forecasts



1. Errors are initially present on all scales

2. Errors in small scales saturate very fast

3. Errors in the synoptic scales grow primarily due to instabilities, not due to inverse cascade process

4. The predictable part of the flow beyond a few hours are mostly in the synoptic scales

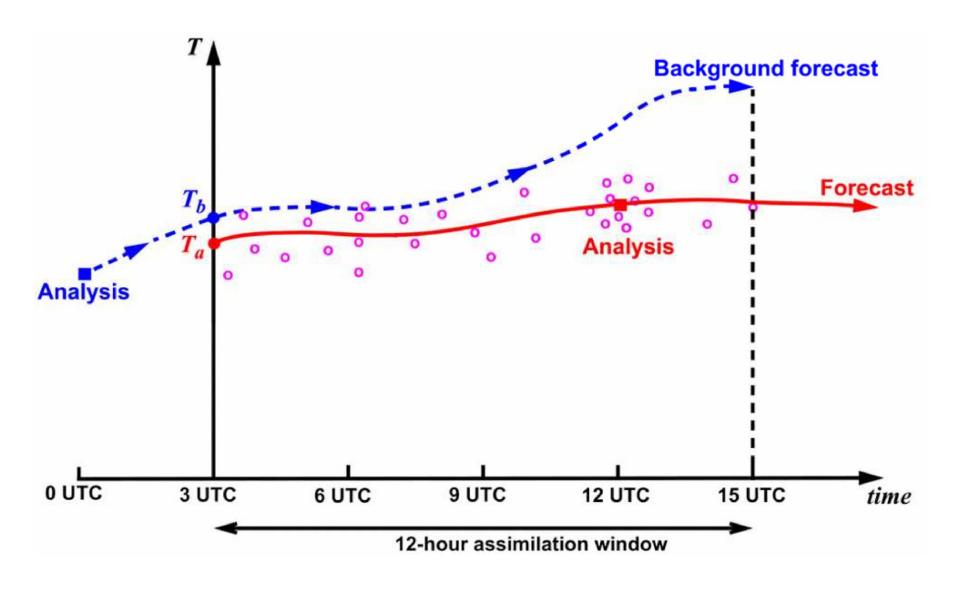
1. Errors are initially present on all scales

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- 3. Errors in the synoptic scales grow primarily due to instabilities, not due to inverse cascade process
- 4. The predictable part of the flow beyond a few hours are mostly in the synoptic scales

Data Assimilation

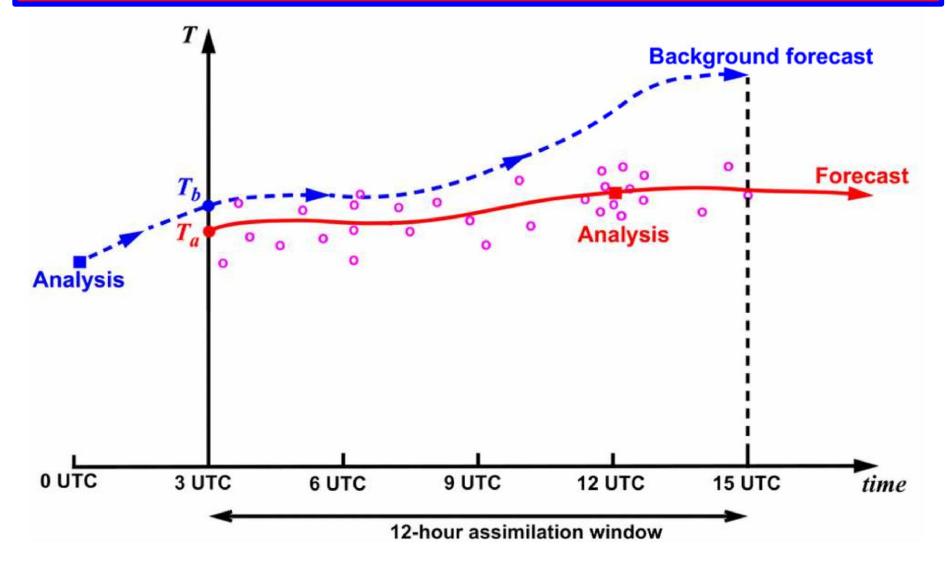
Traditional 4DVAR



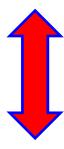
Quantity of available data

Number of Observations = $M \sim 10^5 - 10^6$ Dimension of State Vector = $N \sim 10^7 - 10^9$

$$J(\mathbf{x}_{o}) = J_{b} + J_{o} = \frac{1}{2} \left(\mathbf{x}_{o} - \mathbf{x}_{o}^{b} \right)^{T} \mathbf{B}^{-1} \left(\mathbf{x}_{o} - \mathbf{x}_{o}^{b} \right) + \sum_{i=1}^{I} \frac{1}{2} \left(H_{i} \mathbf{x}_{i} - \mathbf{y}_{i} \right)^{T} \mathbf{R}_{i}^{-1} \left(H_{i} \mathbf{x}_{i} - \mathbf{y}_{i} \right)$$

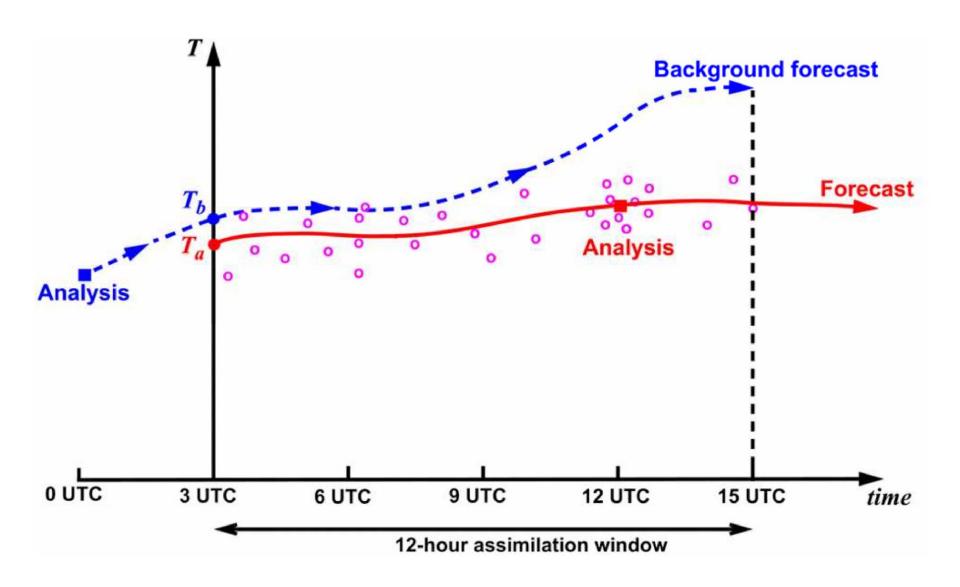


Data Assimilation



Ensembles

$$\mathbf{J}(\mathbf{x}_{o}) = \mathbf{J}_{b} + \mathbf{J}_{o} = \frac{1}{2} \left(\mathbf{x}_{o} - \mathbf{x}_{o}^{b} \right)^{\mathrm{T}} \mathbf{B}^{-1} \left(\mathbf{x}_{o} - \mathbf{x}_{o}^{b} \right) + \sum_{i=1}^{\mathrm{I}} \frac{1}{2} \left(H_{i} \mathbf{x}_{i} - \mathbf{y}_{i} \right)^{\mathrm{T}} \mathbf{R}_{i}^{-1} \left(H_{i} \mathbf{x}_{i} - \mathbf{y}_{i} \right)$$



Quantity of available data

Number of Observations = $M \sim 10^5 - 10^6$ Dimension of State Vector = $N \sim 10^7 - 10^9$

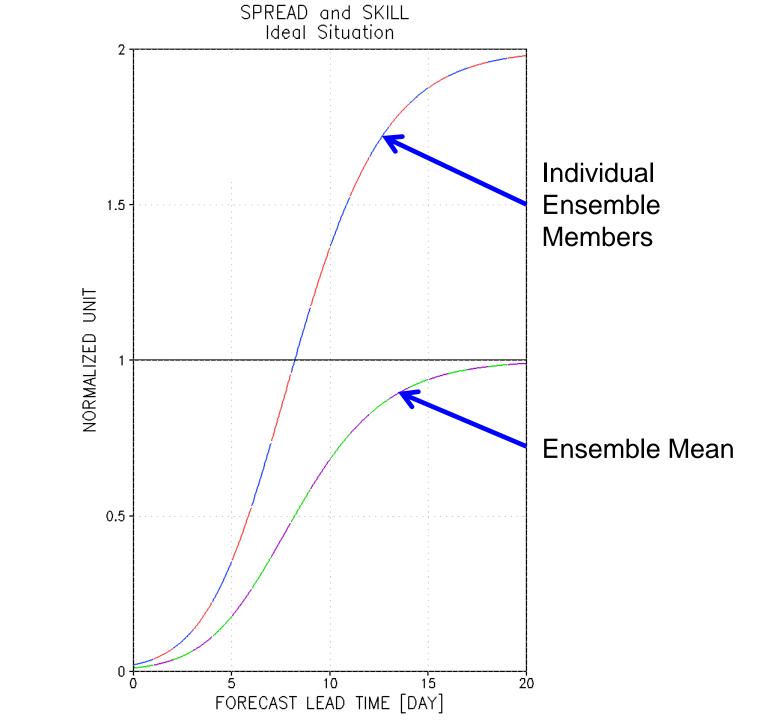
Only the largest scales are really defined by the available data

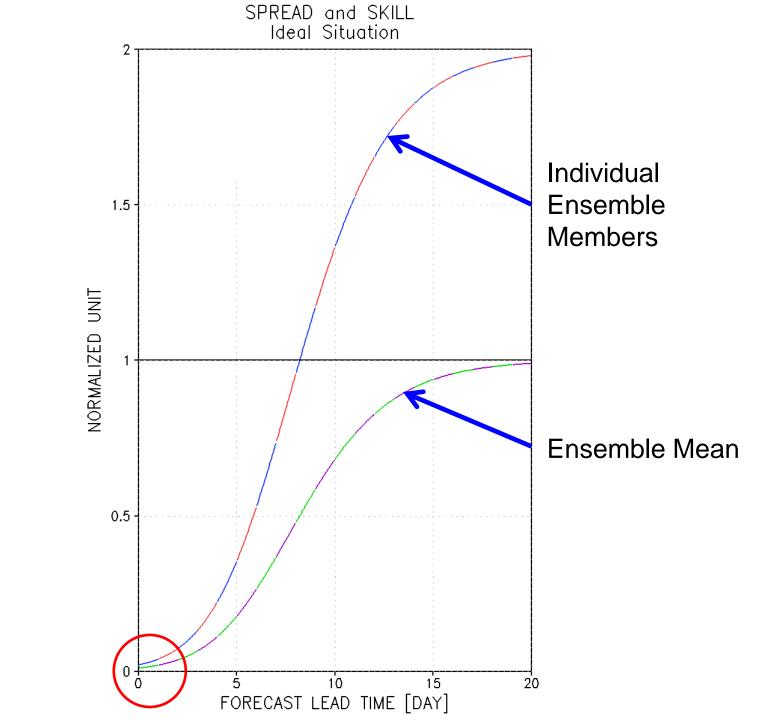
Smaller scales are not

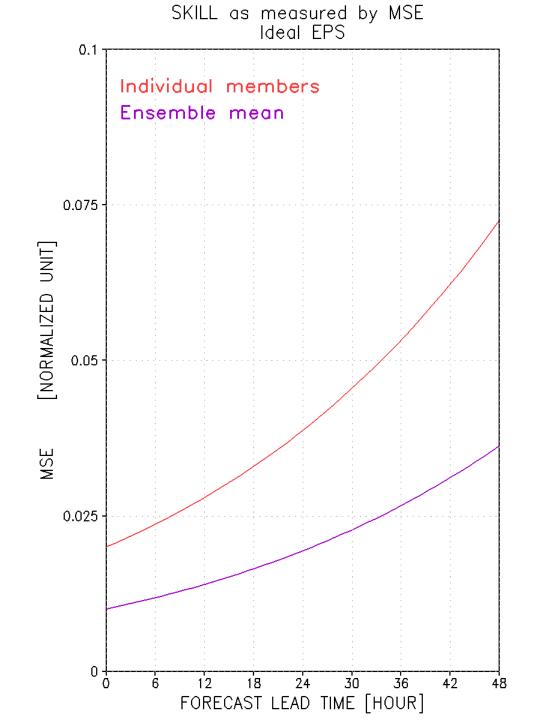
There should thus be an infinite amount of equally likely IC

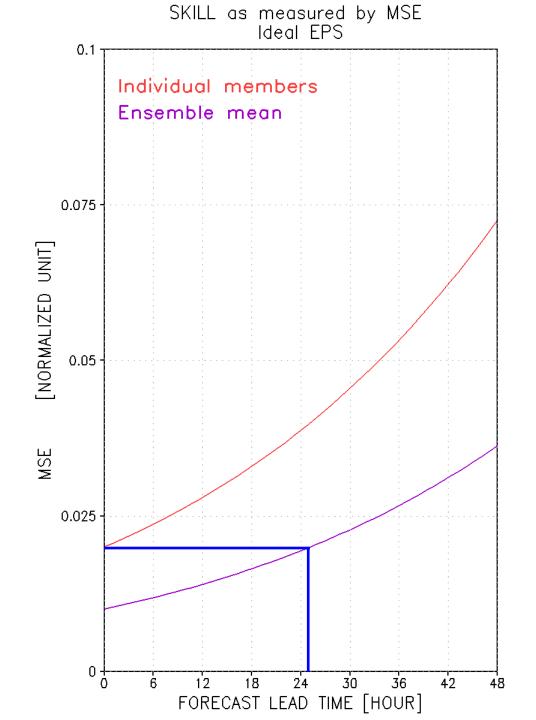


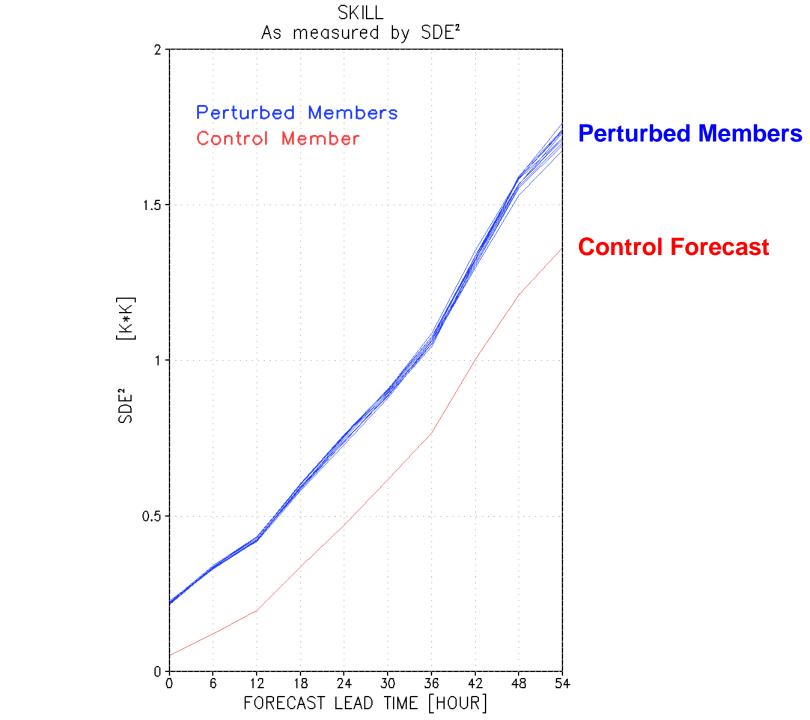
Ensembles of IC

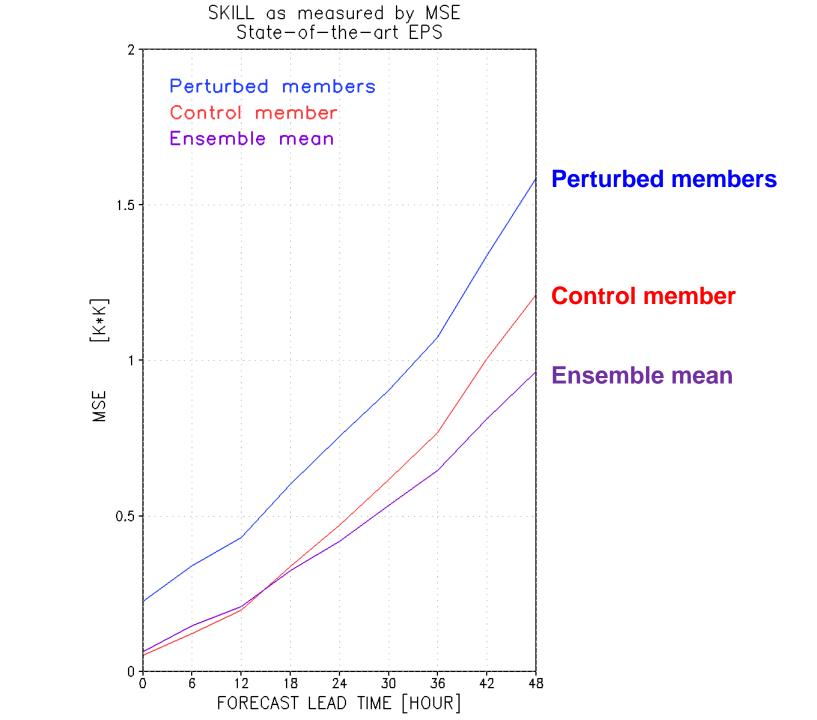




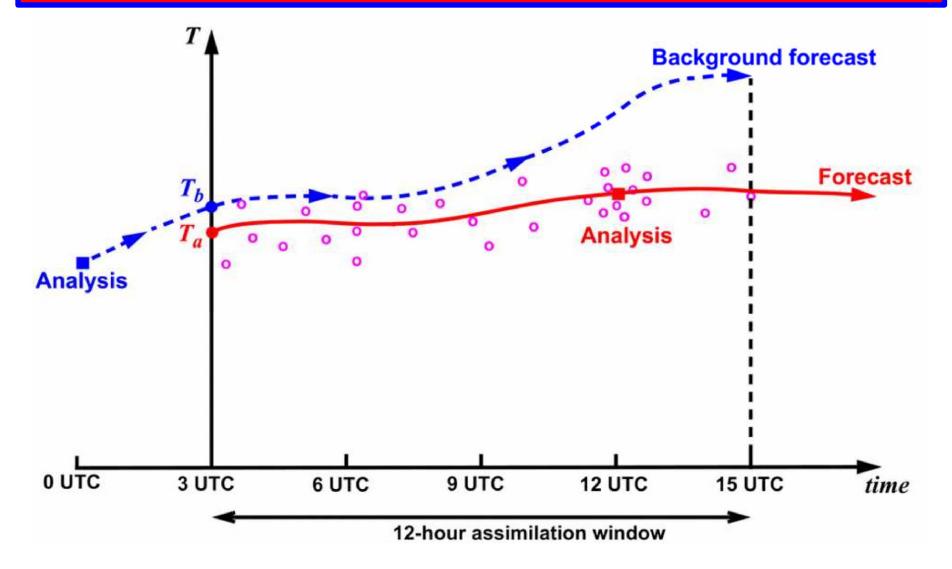


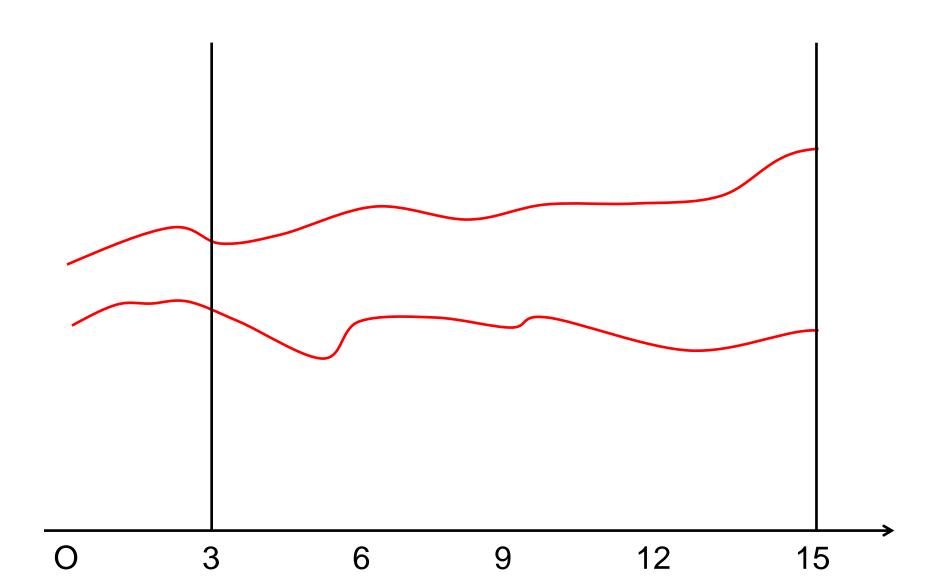


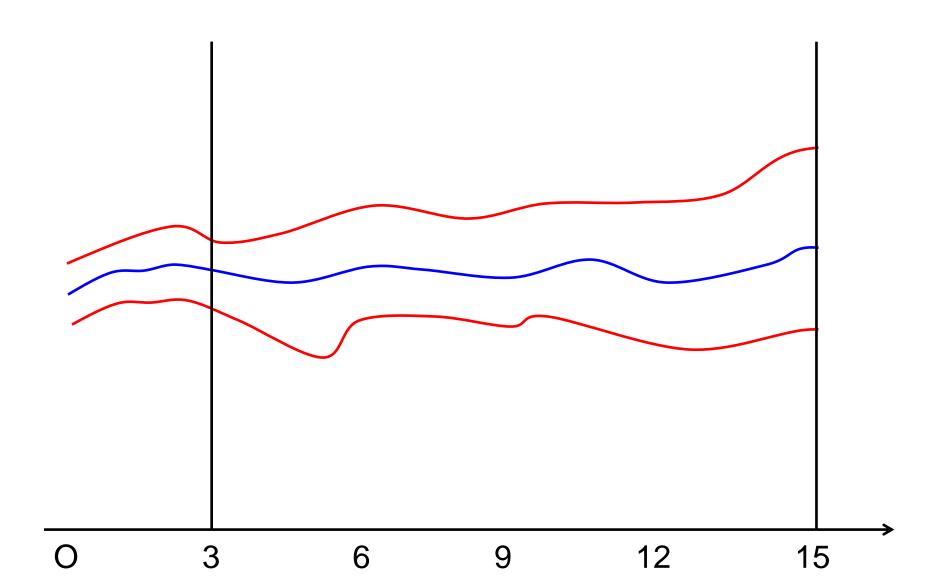


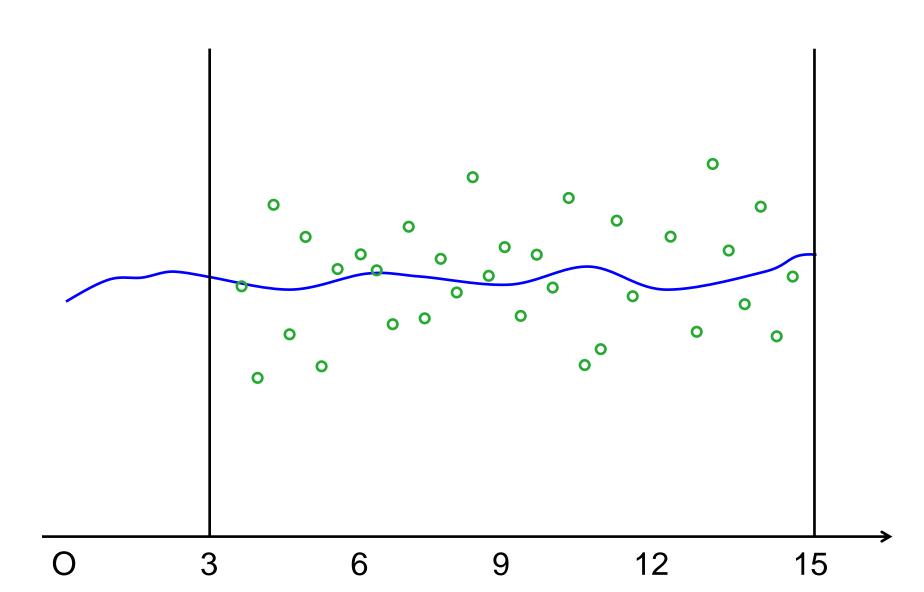


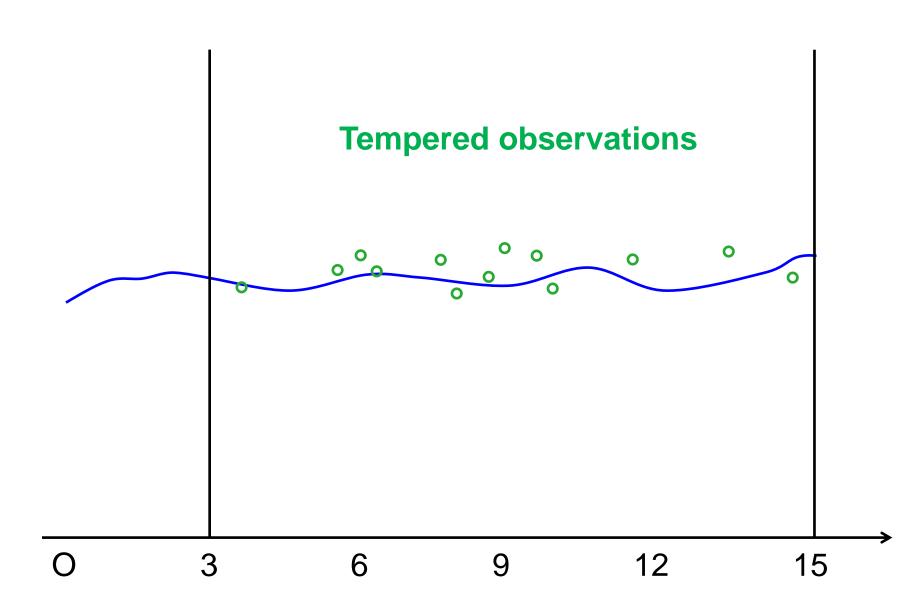
$$\mathbf{J}(\mathbf{x}_{o}) = \mathbf{J}_{b} + \mathbf{J}_{o} = \frac{1}{2} \left(\mathbf{x}_{o} - \mathbf{x}_{o}^{b} \right)^{\mathsf{T}} \mathbf{B}^{-1} \left(\mathbf{x}_{o} - \mathbf{x}_{o}^{b} \right) + \sum_{i=1}^{\mathsf{I}} \frac{1}{2} \left(H_{i} \mathbf{x}_{i} - \mathbf{y}_{i} \right)^{\mathsf{T}} \mathbf{R}_{i}^{-1} \left(H_{i} \mathbf{x}_{i} - \mathbf{y}_{i} \right)$$

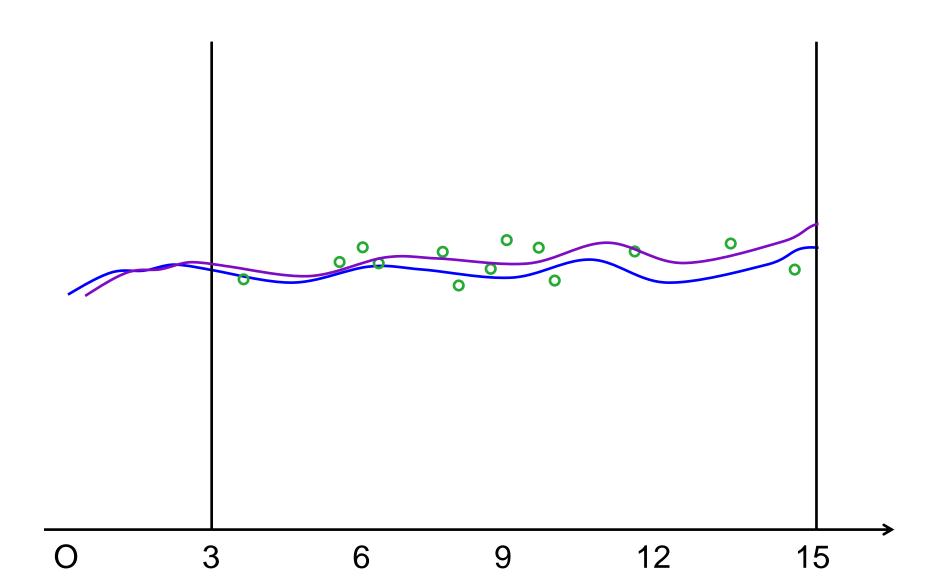


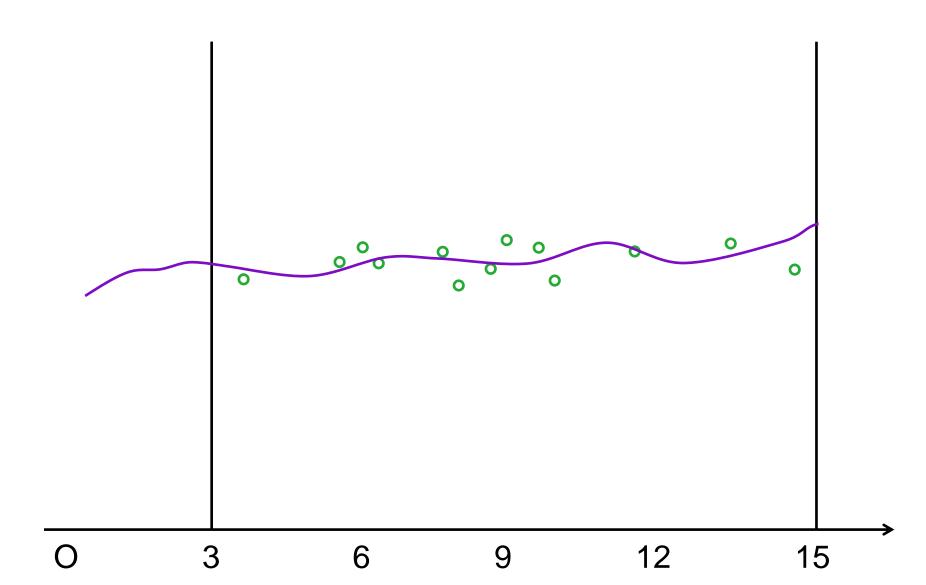


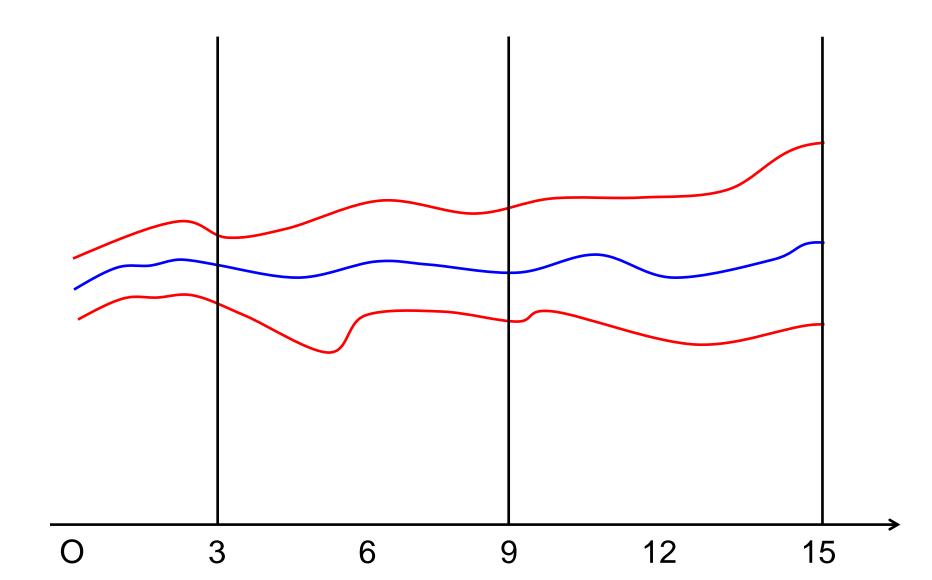


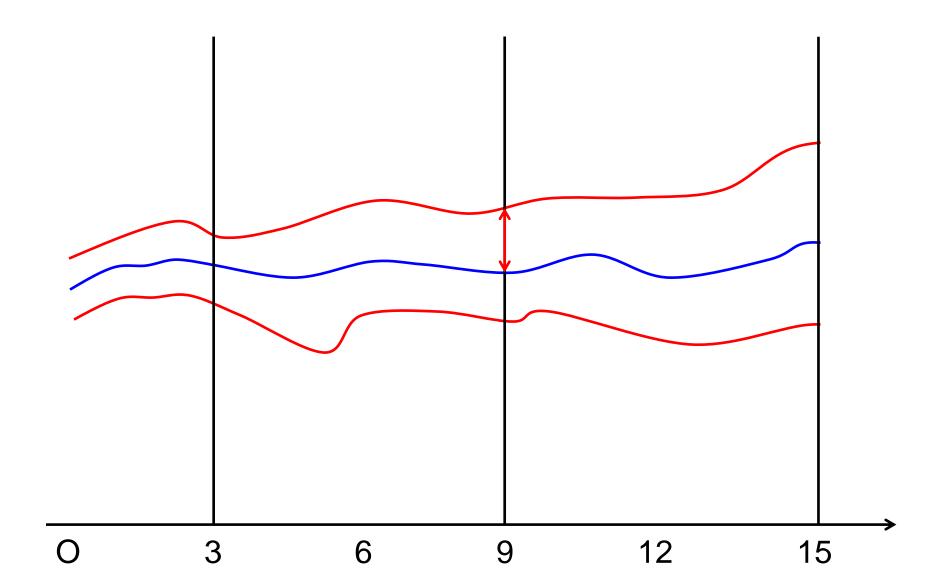


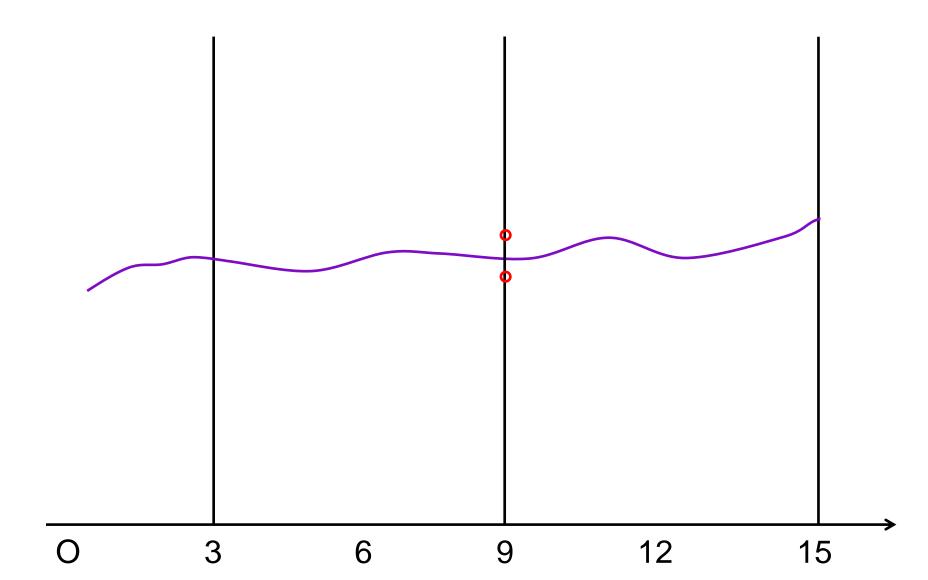


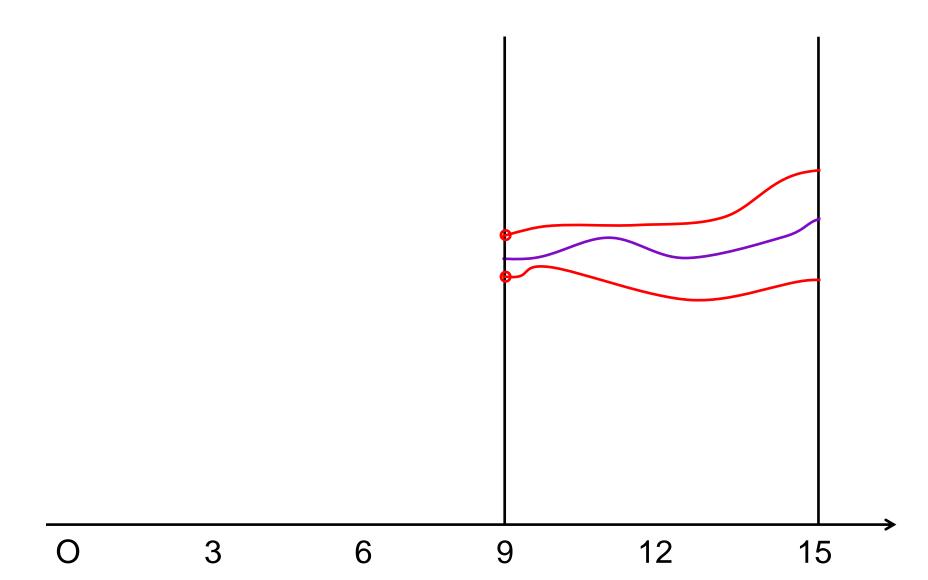




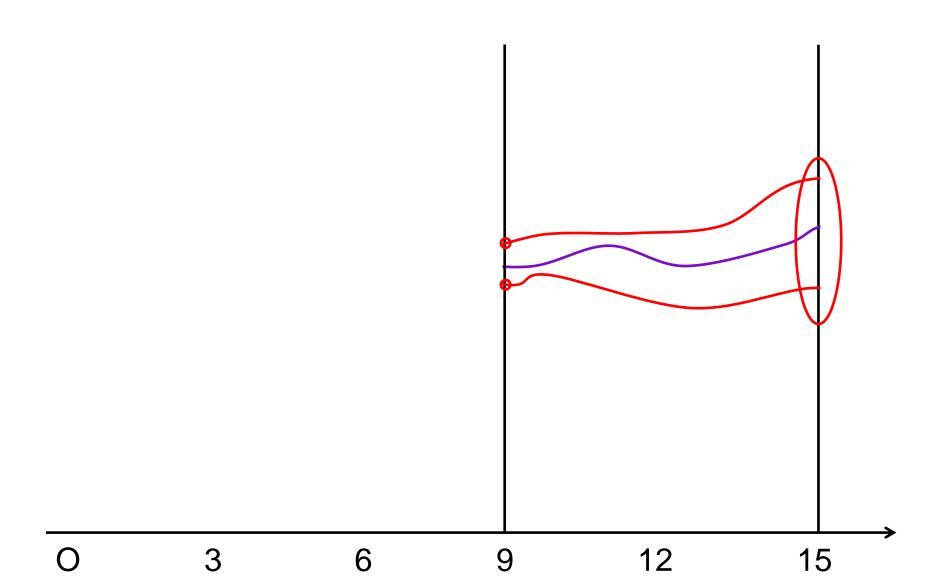




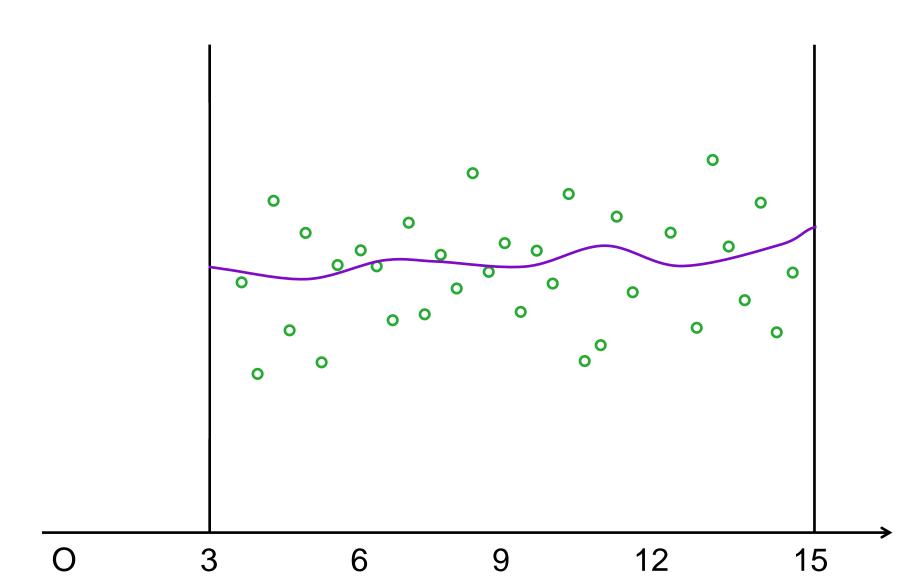




Spread



Skill

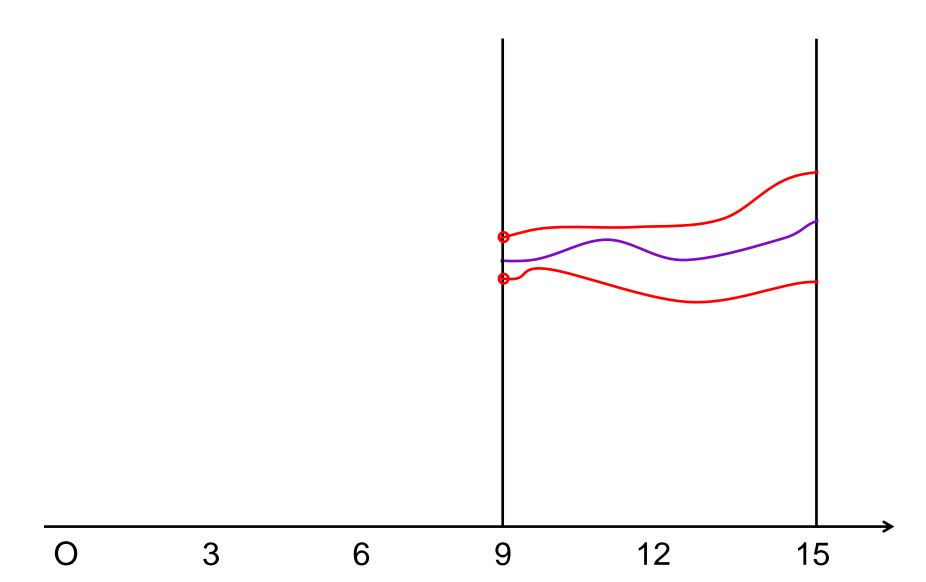


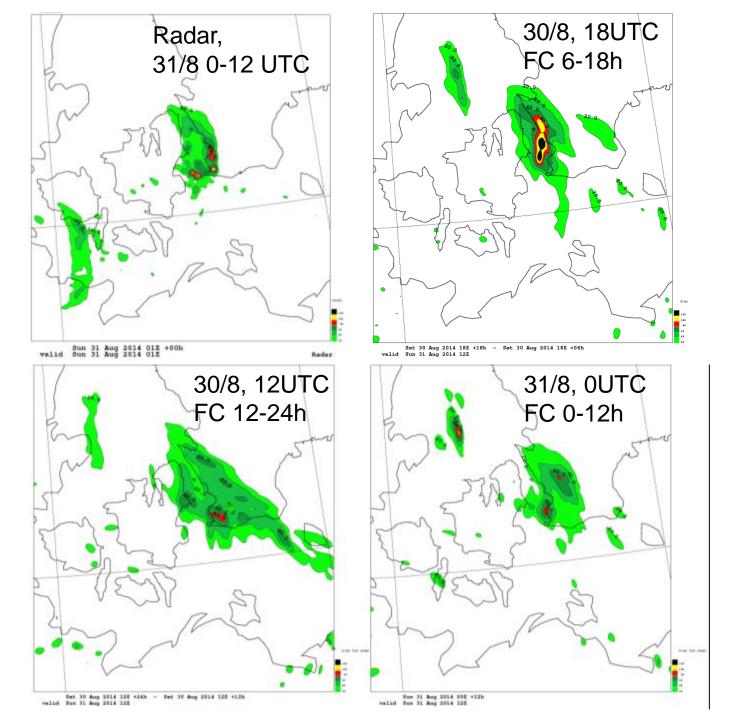
Quantity of available data

Number of Observations = $M \sim 10^{20}$ Dimension of State Vector = $N \sim 10^{15}$

However

Only the largest scales are predictable at the end of the window





Slutsatser

- Synoptiska skalans stora betydelse
- Prediktabiliteten mycket begränsad hos mesoskalan
- Observationsunderlaget långt ifrån tillräckligt att beskriva mesoskalan